

OROGENIC MINERALIZATIONS – A NEW EXPLORATION TARGET FOR GOLD- POLYMETALLIC ORE DEPOSITS IN GREECE

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SDIMI 2013 - 6th International Conference on Sustainable Development in the Minerals Industry, 30 June - 3 July 2013, Milos island, Greece



line Orogenic Gold (OG) setting



A) Accreted terrains. gold deposits orogenic temporally & spatially associated with episodes of crustal growth within developing/growing forearc (to back-arc) region volcanic of arc or continental margin. Contemporaneous with/post-date (medium P-T) metamorphism ot host rocks (epigenetic) Compressional/transpres sional environments in orogenic gold deposits (Goldfarb et al. 2005).





ProMine Mineralization types

Main commodity associations

Number	Association Name	Commodity Association	'Type' codes queried in MD database (including 'sons')	
1	Alkaline & Peralkaline intrusions	Nb, REE, P, (Ta, Zr, Sc, F, U, Fe)	C10, C20	
2	Epithermal	Au, Ag, Sb, Hg, Te, Cu, In	D	
3	Igneous Felsic	Sn, W, Ta, Nb, (Mo, Li, Be, B, In, F)	C40	
4	Igneous Intermediate	Cu, Mo, Au, (Re)	C50	
5	Igneous Replacement	Fe, W, Pb, Zn, Cu, Au	C70	
6	IOCG	Fe, Cu, Au, (P, REE, U, Co)	К	
7	Mafic intrusion	Fe, Ti, V	B30	
8	Mafic or UltraMafic	Ni, Cr, Cu, PGE, (Co, Bi, U, Ag)	B, except B30	
9	Orogenic Gold	Au, (Ag, As, W, Cu, Sb, Bi)	A, plus 'commodity Gold'	
10	Pegmatites	Nb, Ta, Sn, Li, Be, (U, REE)	C60	
11	Carbonate-hosted deposits	Zn, Pb, Ag, Ba	F40	
12	Sandstone- and shale- hosted deposits	Cu, U, Pb, (Ni, Co, Zn, V, PGE, Re)	F20, F30, F60	
13	Sedimentary deposits	Fe, Mn, Ba,K,Na,Sr	F50	
14	1440	Cu, Zn, Pb, (Ag, Au, Te, Sn,	_	
	VMS	In)	E	
15	Residual deposits	In) Fe, Al, Ni, Cu, (Mn, Au, P, REE)	E H20, H30	

In order to analyse homogeneous mineralization, a list of 16 most characteristic commodity associations – or deposit types - has been established by ProMine partners.

The ProMine MD database was then queried to extract all deposits of these 16 major types.

As a result, 16 homogeneous deposit populations were obtained, that were processed for potential and predictive mapping.





ProMine Mineral potential

Mineral potential mapping - results





ProMine Pan-European approach

Mineral potential mapping - results



Orogenic gold:

Distribution of potential is guided by a single main commodity (Au) and a well constrained type of mineralization. Major districts belong to two groups:

-Paleoproterozoic orogenic deposits related to greenstones in the Fennoscandian shield;

-Hercynian gold-bearing districts related to late Hercynian (~300 Ma) deformation belts (N. Iberian peninsula, French Massif Central, Bohemian Massif).

Additional more scattered deposits can be found in other Hercynian (Salsigne) or Caledonian (Great Britain and Norway) domains and the Balkan-Carpathian region.





Orogenic gold deposits are a distinctive type of mineral deposits that has been the source for substantial world gold production.

- The globally youngest dated orogenic gold lodes are the very widespread vein systems of the Alpine Carpathian orogen which are associated with Europe / Africa collision during the Palaeogene.
- The Alpine Carpathian orogen underwent thrusting, nappe emplacement and high temperature metamorphism in the latest Eocene, followed by Oligocene vein formation.
- The best studied orogenic gold deposits in the Alpine Carpathian orogen are the deposits of the Monte Rosa province in northwestern Italy, the orogenic gold veins in to the Austrian Alps and the Late Miocene veining occurred in the Swiss Alps at about 10 Ma.





ing Subduction-related magmatism











Major Au mineral zones in Greece







Greek OG setting & styles

Orogenic gold deposits

- are present in various metamorphic terrains displaying various types of mineralization.
- show strong structural control of the ore forming processes and potential deposition takes place near large scale compressional structures (thrust shear faults).
- their controlling structures are mainly related to major faults and associated brittle deformation characterized by silicified fracture zones, ductile shear zones and breccias, and foliated zones.

Three styles of orogenic gold-base metal mineralizations have been defined in relation and association with

- Iow-grade metamorphic rocks.
- overthusted serpentinites.
- high-grade metamorphic terrains.





Adding a new Au target





OG in low grade metamorphic rocks

Komaros area



 Chlorite, chlorite-mica schists,phyllite series. 2.Calc-schists. 3.Metaconglomerates. 4.Metasedimentary rocks. 5.Basal conglomerates. 6.Clastic sediments. 7.Nummulitic limestones. 8.Altered epiclastites and volcanoclastic sandsiones. 9.Andesites (aglomerate tuffs and lavas). 10.Fracture zones. 11.Epithermal zones and veining. 12.Thrusting (base metals and gold mineralization).

Geological map of the Kommaros and Perama gold deposit (Michael C., 2002)



Structural setting

The mineralization in eastern Rhodope zone is associated with thrusting and hosted by brecciated greenschists and calc – schists (Circum Rhodope belt). The thrust follows the contact between greenschists and calcschists.

The structural pattern observed in the area is interpreted to have formed under north – south compressional conditions





oMine OG in low grade metamorphic rocks

Mineralization in folds and veins.



Mineralization occurs in quartz veins and folds (en-echelon folds, convergent and conjugate folds). Au, Ag, Cu, Pb, Zn, As is the characteristic geochemical association.

 Albite, carbonate minerals, chlorite, sericite, epidote and barite are common gangue minerals





Mine OG in low grade metamorphic rocks

Oxidized ore as matrix of brecciated material



Ore mineralogy and grades

- The mineralization is characterized by oxidation of base metal sulphide ore minerals. Gold occurs as native gold in quartz veins and as inclusions in iron oxides. The oxidized ore hosts gold from 1 to 49 g/t and the average grade is 6,3 g/t. Galena and sphalerite are the residual primary minerals in oxidized ore.
- The oxidized ore consists of smithsonite, cerussite, malachite, azurite, hematite, jarosite and gold. The proximal gangue mineral assemblages consist of albite, carbonate minerals, chlorite, sericite and epidote, minerals typical for orogenic gold mineralization





- This is well developed in western Rhodope zone.
- The deposits are normally controlled by second order structures near the compressional thrust fault between Serbomacedonian and Rhodope zones.
- The common characteristic of these mineralizations is mainly the Fe-rich oxide mineral gossaneous concentrations.
- The host rocks are varying with marbles and schists to be the most common ones.



OG in high metamorphic terrains

•Kalindiri Angistro Le?

Structural setting

- Two main compressional deformation events (D1, D2) were recognized in crystalline units of the western Rhodope zone. (D1), of
 Eocene-Oligocene age (50-23 Ma), created a system of NE-SW trending folds and associated reverse faults.
 (D2), of Miocene age (18-23 Ma), is accompanied by the Kavala granodiorite intrusive (19-22 Ma).
- The deposits are normally controlled by second order structures near the compressional thrust fault between Serbomacedonian and Rhodope zones.



OG in high metamorphic terrains



In Angistron the mineralization has low volume and is associated with a zone of hydrothermal alteration several hundred of meters wide. The alteration is also related to a series of subvertical vein systems.

- The mineralizations in western Rhodope zone are located in the Pangeon, Palea Kavala, Symvolon and Angistron.
 - Mineralization in Palea Kavala is a combination of subhorizontal thrust shear zone and subvertical fault. The intersection of these structures has provided a location favoring dissolution and replacement of the marble. Large auriferous pods have been developed.
- Mineralization at Symvolon is controlled by thrust which occurs along the contact between the upper marble horizon and lower schist or gneiss zone.





OG in high metamorphic terrains

Au – Sb polymetallic
 mineralization in Kalindiri area.



•Intensive alteration zones associated with faults (normal and thrust faults) in marbles characterized by quartz, dolomite, calcite kaolinite, ankerite, garnierite, jasperoids.

Structural setting and mineralogy

- formed under the influence of the major regional north – south compressional environment. Thrust faults trending eastward are very important structures for the area.
 - Au Sb polymetallic mineralization is normally controlled by second order structures near the compressional thrust faults. The NW, NNW trending shear zones are the most important controlling structures. The thickness of the fault zones is up to five meters and the zones are strongly brecciated.

found in silicified marbles along the thrust fault contact between carbonate rocks and metamorphic basement rocks Disseminated free gold occurs in a quartz matrix. Gold (5-23 μ) also occurs within Fe – hydroxides.



Microthermometric analyses

	Type of mineralization	Mineral	Ran ge (Th °C)	S a l i n i t	CO2 content	Host rocks		Location	Source of data	
				y						
y d	Antimony- polymetallic veins	Quartz	294- 330° 187- 261°		CO ₂ - H ₂ O rich fluids	Gneisses		Kalindiri	personal ication)	
	Gold mineralization in silicified thrust zones	Quartz	150- 268°			Marbles	Kalindiri		Kilias S. (commun	
	Gold associated with serpentinites	Quartz	250- 350°			ntes	Esochi -Akritas		Iliadis, 2005; Kilias S.	
		Sphalerite (216- 244º			Serp entir			(personal communication)	
	Shear zones controlled gold mineralizations veins	Quart z		CO ₂ - H ₂ O rich fluids	es, Schists		vala	Naden J 1994; Nimfopoulos,		
		Calcite	320°		H ₂ O rich fluids	Marbl Gneisses,		Palia Ka	2000; Eliopoulos 1993	
1000	Shear zones controlled gold mineralizations	Quartz	275- 330°		CO ₂ - H ₂ O rich fluids	Marbles, Gneisses, Schists	Gneisses, Schists Pangeon		Naden S. 1994; Nimfopoulos, 2000	
		Quartz	216- 420°							
	Sheeted quartz vein system	Quartz	216- 420°		CO ₂ - H ₂ O rich fluids	Kavala		Kavala	,2011	
		Quartz	210, 4- 324, 4°		CO ₂ - H ₂ O rich fluids				Fornade	
	Sheeted quartz vein system	Quartz	255, 7-		CO ₂ - H ₂ O rich			alkero		

ð

414°

fluids

The homogenization temperatures (Th) clearl indicate hydrothermal fluids typical for orogenic gol mineralizations and deposits. For example:

 the Sb mineralization in Kalindiri area ranges from 294° to 330° and the salinity is 0,2 - 7 wt% NaCl equiv. The homogenization temperatures suggest also that the mineralization took place during the gradual decrease of fluid temperature from 330° to 118°. Gold mineralization was deposited between 118° and 268°C. The coexistence of vapor rich and liquid rich inclusions is may be an evidence of boiling.

•the Au – bearing veins of Palia Kavala and Pangeon areas indicate that mineralization took place between 216 and 440° (low salinity, H₂O-CO₂ rich fluids), conditions similar to orogenic environment of deposition

•the large quantities of CO₂ and H₂O-rich fluids were released during greenschists to lower amphibolites facies of metamorphism in the Rhodope zone





Stable isotopes (O,C) analyses

	δ ¹⁸ Ο	Mean	Num. of	δ ¹³ C	Num. of	Source
Area	(smow)	value	samples	(PDB)	samples	of data
	(Range)					
Palea	+14,37	+22,01	Area	-9,19	13	Baker,
Kavala	to			to		1991
8 11	+26,55		and a start	+2,11		
Pangeon	+26,11	+28,72	15			Baker,
	to					1991
	+30,92					

Stable isotopes (O,C) analyses of calcite

The oxygen isotopes analyses of calcite veins from Palea Kavala and Pangeon with average δ^{18} O values 22,01 and 28,72 per mill respectively, support the metamorphic origin (+15 to +35 per mill) of the fluids.





Conclusions

Gold and base metal mineralizations occur throughout the Rhodope zone in northern Greece at a range of spatial scales and geological environments. Orogenic gold deposits in Greece are present in various metamorphic terrains displaying variable types of mineralization. The main geological characteristics include:

- Deformed and variably metamorphosed host rocks
- •Strong structural control of the ore forming processes and potential deposition taking place within/near large/small scale compressional structures (thrust faults)
- •The emplacement of granitoid stocks involves increase in heat flow and fluid circulation in response to fracturing.





Conclusions

•The fluid inclusion data, the stable isotopes values and some geological data define a geochemical environment indicating orogenic gold mineralization setting.

•The mesothermal conditions of deposition, the low salinity, the CO_2 - H_2O rich fluids, the high oxygen values are characteristic for orogenic type of mineralization.

•The emplacement of the Kavala granitoid took place 19-22 Ma,) and coincides with the second (D2) compressional deformation event (18-23 Ma). This may indicate, the orogenic gold veins occurring in the western Rhodope zone belongs to Miocene. Absolute dating of the above orogenic mineralizations is still open to further investigation. The detailed knowledge of the orogenic mineralizations and processes of deposition can help to define potentially mineralized areas within regions of Rhodope and Serbomacedonin zones.

•This is the first documentation of orogenic mineralizations which become a new exploration target for gold-polymetallic ore deposits in Greece.





OROGENIC GOLD PREDICTIVITY

Derived from mineral deposit density, weighted by deposit class





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Thank you for your attention



